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Adnand Bitri, Kevin Samyn, Stéphane Brulé, Emmanuel Javelaud. Assessment of ground compaction using multi-channel analysis of surface wave data and cone penetration tests. *Near Surface Geophysics*, 2013, 11 (6), pp.683-690. 10.3997/1873-0604.2013037. hal-01025230

HAL Id: hal-01025230

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Submitted on 17 Jul 2014

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Assessment of ground compaction using multi-channel analysis of surface wave data and cone penetration tests

A. Bitri^{1*}, K. Samyn¹, S. Brûlé² and E.H. Javelaud²

¹BRGM, Bureau de Recherches Géologiques et Minières, Risks Department, 3 Avenue Claude Guillemin BP36009 45060 Orléans Cedex 2, France

²Ménard, 2 rue Gutenberg, BP28, 91620 Nozay, France

Received March 2012, revision accepted May 2013

ABSTRACT

The construction of a new industrial and commercial area in 2009 at the Givors' former glass factory area in France involved heavy dynamic compaction work. For the purpose of founding the new buildings, it was necessary to improve the ground mechanical properties of 7–15 m of well-graded gravel backfill lying on geotechnical bedrock. In order to assess the quality and depth of ground compaction, cone penetration tests are often performed before and after compaction. The method is intrusive and a one-location test. It requires a substantial amount of time to evaluate a large area and evaluation quality is quite dependent on the operation technique and soil type. In this paper, the quality and extent of ground compaction were evaluated using results from the *in situ* Multi-Channel Analysis of Surface Waves (MASW) seismic method and cone penetration tests (CPT). MASW tests were used to determine shear-wave velocity (V_s) profiles before and after compaction and CPT tests were adopted to determine the correlation between V_s and the measured penetration resistance (q_c) improvement along profiles. The results of this study show the effectiveness of surface waves for the evaluation of compaction performance and demonstrate the potential of this technique to engineering and environmental problems.

INTRODUCTION

Evaluation of soil geotechnical parameters is a preliminary task to be conducted either for designing building foundations, or for studying site effects in seismic hazard evaluation. Both of these applications show the importance of knowing soil elastic properties, which govern its behaviour in presence of natural or human solicitations. Shear-wave velocity (V_s) is usually taken as a good indicator of the elastic behaviour of soils and is classically estimated from cross-hole measurements. In order to avoid destructive and expensive testing, we study here the possibility to estimate the quality and depth of ground compaction using V_s distribution with depth obtained by the Multi-Channel Analysis of Surface Waves (MASW) method (Park *et al.* 1999). This method studies surface wave dispersion i.e., variation of phase velocity (c) with frequency (f). The technique consists in: 1) data acquisition using a multichannel recording system; 2) dispersion curve picking; 3) inversion of dispersion curve. Surface waves are easy to generate by impact sources or vibrators and are mostly sensitive to shear-wave velocity. The determination of V_s profiles from surface waves is attractive since their propagation velocity, i.e., phase velocity, highly depends on the stiffness-depth profile. An important step in the MASW method is to generate a reliable

image of dispersion energy in the frequency velocity (f - c) domain. A high-resolution image leads to an accurate picking of dispersion curves and helps to separate fundamental and higher propagation modes of surface waves. Several methods allow to obtain dispersion images such as the F - K transform (Gabriels *et al.* 1987), τ - p transform (McMechan and Yeldin 1981; Moktar *et al.* 1988), phase shift method (Park *et al.* 1998) and linear Radon transform (Luo *et al.* 2008). Modelling results (Park *et al.* 1998) show that the resolution of a dispersion image is mainly influenced by the geophone spread length. The resolution increases as the geophone spread increases.

Estimating V_s variations with depth is tackled as an inverse problem for which the phase velocity variations with frequency constitute the observations. Several inverse methods used in seismology have already been proposed to resolve such a problem. Dorman and Ewing (1962) used surface waves to determine elastic properties of crustal-mantle structures. Engineering problems were tackled in Nazarian and Stokoe (1984) where a hammer was used as an impulsive source for generating Rayleigh waves while ground motions were recorded from two seismic sensors. The dispersion was afterwards derived from the spectrum of recorded vertical motions. Addo and Robertson (1992) also used the same approach and validated the method on six different sites by comparing Spectral Analysis of Surface Wave

* bitri@brgm.fr



FIGURE 1

a) Photograph of the dismantled Givors' former glass factory site, purpose of the dynamic compaction project; b) conceptual 3D view of the new industrial and commercial area. (from Givors Development, by courtesy)

(SASW) (Nazarian and Stokoe 1984) inverted results and cone penetration test (CPT) data. The evaluation of ground densification using SASW and resonant column (RC) tests was performed by Kim and Park (1999). Gabriels *et al.* (1987) and Karray *et al.* (2010) improved the method by taking into account higher propagation modes of Rayleigh waves and by increasing the distance between the sources and receivers. Xia *et al.* (1999a) proposed an inverse method based on the Levenberg-Marquart technique. Recently, inversion of the full-waveform has been proposed in order to overcome the limitations of the usual approach based on the inversion of a dispersion curve (Forbriger 2003a,b; Romdham *et al.* 2011.). The Linearized Least-Squares (LLS) technique used for this study is adapted from Hermann (1987) and was already tested in Bitri *et al.* (1998).

The purpose of this paper is to present and discuss the results of our methodology based on a surface wave method for the assessment of the quality and depth of ground compaction on the studied site.

SITE DESCRIPTION

The project is located on the Givors' former glass factory area along the Gier river. The site was occupied for over a century by industrial activities. The plant was dismantled recently (Fig. 1a) and the construction of new industrial and commercial buildings is planned (Fig. 1b). For the purpose of building and founding these structures, a ground improvement technique was proposed by engineers to meet the dual problem of soils with low-mechanical properties and environmental issues resulting from a chemical residual impact of former industrial activities on the soil and water. A compaction program was developed in order to improve the backfills overlying sand and gravel deposited by the Gier river. Particle size analysis of five samples taken at 2 m depth and average friction ratio (f) clarified that the filling material is of sand type. According to geotechnical available reports, the average mechanical characteristics of the backfills are: $q_c = 3.7$ MPa, $f = f_s / q_c \sim 1\%$, $pl^* = 0.53$ MPa and $E_M = 8.2$ Mpa where q_c is the measured cone resistance obtained during CPT tests, f_s the friction ratio obtained during CPT tests and pl^* and E_M are breaking limit pressures. For the purpose of assessing the quality and depth of the ground compaction procedure, preliminary tests were realized over a 160 x 35 m area (Figs 1 and 2). The positions of the four parallel seismic lines (A, B, C and D) and the five CPTs (B3, B15, C8, D3 and D15)

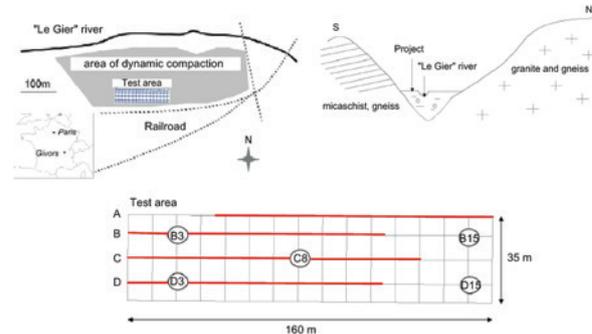


FIGURE 2

Localization of the 160 x 35 m test area on the dismantled Givors' former glass factory site. The four parallel seismic lines (A, B, C and D) are highlighted in red solid lines and the five CPTs (B3, B15, C8, D3 and D15) are marked using black circles. (Adapted from Brûlé *et al.* 2010b)

that were performed before and after dynamic compaction on this test area are shown in Fig. 2. The groundwater level is about 6 m below the surface. Initial investigations and geotechnical feasibility studies have recommended pile or micropile foundations, embedded in compact sand and gravel at 12–15 m depth. In fact, deep piles could impact the water resource, blending the superficial polluted soil and water with the deeper ones.

DYNAMIC COMPACTION

Dynamic compaction is a cost attractive ground improvement technique for brown sites especially when thick, heterogeneous and polluted backfills have to be densified.

'High Energy' Dynamic Compaction (HDCTM) work was performed in 2009 at the Givors' former glass factory area in France. For the purpose of founding the new buildings, the objective was to improve the ground mechanical properties of 7–15 m of well-graded gravel backfills laying on geotechnical bedrock composed of sediments of the Gier river.

For granular soils, dynamic compaction tends towards an improved soil with more homogeneous properties than initially. Therefore, for this 85 000 m² building project, the heavy compaction work increased the density of the ground by means of energy applied by ponder ($m = 27\ 000$ Kg) free-dropped on the soil surface from various heights ($H_{dro} = 20$ –25 m) with a square print grid (Fig. 3). It led to a large-scale ground improvement.

The energy brought at each drop is more than 4000 kJ and the velocity of the ponder impacting the soil is about 20–25 m/s. Ponder drops create compaction craters (Fig. 3) that are filled using the same material before being compacted again. This way, there is no change in the lithology. At the final stage of compaction, the elevation of the area is globally changed by about a –50 to –80 cm downwards vertical translation. In the rest of the study, we assume that this difference in elevation can be considered negligible.

SHEAR-WAVE (V_s) PROFILING

MASW is an effective tool for obtaining vertical shear-wave profiles from a single non-invasive measurement. The method was developed by the Kansas Geological Survey and can be used with active or passive sources (Song *et al.* 1989; Xia *et al.* 1999a; Park *et al.* 1999; Park *et al.* 2007).

Numerical relationships between soil mechanical properties and V_s have been previously published. One of these relationships, which is given by the elastic theory and is an essential property for evaluating dynamic responses and the stiffness of soil, is the small-strain shear modulus, G . V_s is a valuable indicator of the dynamic properties of soil and rock because of its relation with G , given by the equation $G = \rho V_s^2$, where soil density ρ is the total unit weight of the soil divided by gravity. G has units force per length squared (i.e., kPa). This relation demonstrates the good correlation of V_s to soil mechanical properties, which are improved during compaction stages and allow for the expectation of a V_s increase after ground compaction.

The MASW method is based on dispersive characteristics of surface waves. The accepted rule of thumb for the maximum penetration depth in dispersive media is approximately half the longest wavelength. In layered media, the velocity propagation of surface waves depends on the frequency (or wavelength) of the wave because waves of different wavelengths sample different parts of the layered medium. For example, high-frequency (short wavelength) waves propagate only in near-surface layers.

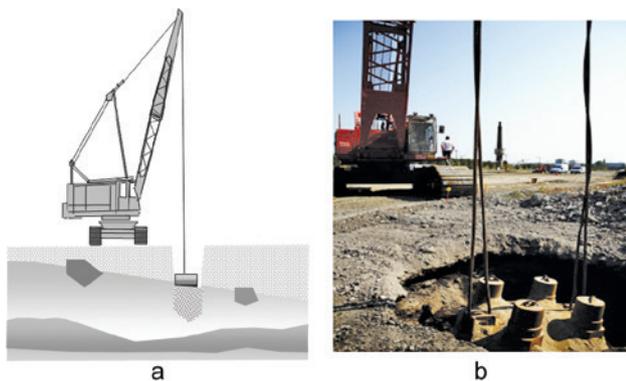


FIGURE 3 a) Schematic representation of the ‘High Energy’ dynamic compaction (HDCTM) procedure; b) photograph of a 27 000 Kg ponder free-dropped on the soil surface.

Low-frequency waves with longer wavelengths propagate through the near-surface layers and also deeper layers. Therefore, by using surface waves over a wide range of frequencies, one can effectively sample different portions of the material profile (Richart *et al.* 1970).

V_s reference models before compaction were established in September 2009 along four linear parallel lines (A, B, C and D) on the test area (Fig. 2). The second surface wave investigations were carried out at the end of October 2009, after heavy dynamic compaction on the same lines in order to assess the effect of compaction.

To increase the speed and efficiency of data recording and thereby keep acquisition costs down, a multichannel seismic cable was designed and manufactured. It consists in 24 geophones at fixed 2 m intervals. Each geophone is attached to a single self-orientating, gimbals-mounted, vertical geophone with 10 Hz central frequency. To help ensure proper coupling, each gimbal geophone is housed in a heavy casing (~1 kg). To damp the motion of the sensor around its rotational axis, the inside of the casing is filled with viscous oil. The seismic cable is towed behind a vehicle. A 24-channel Geometrics Stratavizor seismograph was used to record impacts of a 10 kg hammer seismic source. The source-to-nearest-receiver offset was 4 m, whilst the increment between source locations along the seismic lines was 10 m. The general configuration of receivers and sources in the MASW test is shown in Fig. 4. The record length was selected as 1024 ms at 1 ms sample intervals.

The most critical part of surface wave analysis is the accurate estimation of dispersion curves (i.e., phase velocity as a function of frequency) and surface wave mode. The dispersion patterns were imaged with the slant-stack method in a common shot gather, followed by a 1D Fourier transform over the intercept time (McMechan and Yeldin 1981; Moktar *et al.* 1988). The transformation procedure can be described by the following expression:

$$U(c, f) = \sum_{i=1}^N B^{-1}(f) A(x_i, f) e^{j\varphi_i} e^{-\frac{2\pi i f x_i}{c}} \tag{1}$$

where c is the phase velocity, N is the number of traces in the shot gather, $B(f)$ is the amplitude spectrum of the first trace, $A(x_i, f)$ is the amplitude spectrum of the trace at distance x_i and φ_i is its phase spectrum. The dispersion curve is obtained by picking

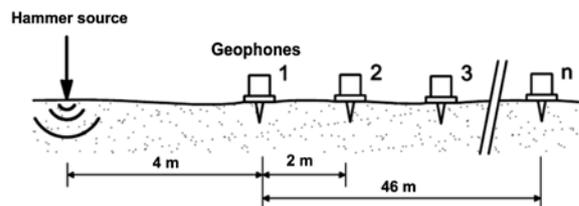


FIGURE 4 MASW test field configuration.

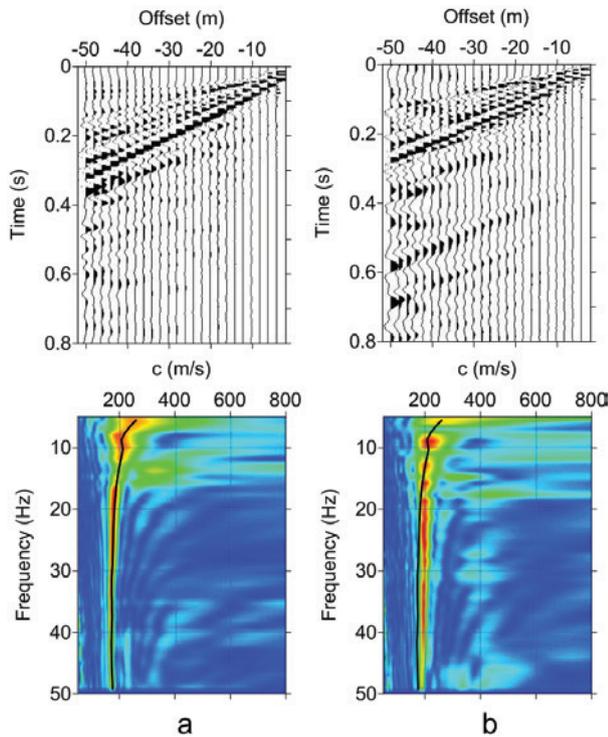


FIGURE 5

Typical seismic shot gathers of line C before (a) and after (b) compaction, with the related dispersion f - c images. The reference fundamental mode before compaction overlays both f - c images in black solid lines. Here we observe the c shift for high frequencies (from 15 Hz) between fundamental modes before and after dynamic compaction.

the maximal values of the $U(c,f)$ modulus. If the seismic signal in the shot gather consists in single noise free surface wave mode, the maximum value of the $U(c,f)$ modulus should be equal to N . If this value is smoother than N , than we can attribute this to an error in the phase velocity between the traces in the shot gather (Hermann 1987).

O'Neill (2003) evaluated, by numerical and field repeatability testing, the error envelopes of dispersion curves determined using equation (1). The observation was a large increase in the phase velocity standard deviation at low frequency, often exceeding 30%, although at the highest frequencies it is 1% or less. In order to qualitatively assess the reliability of the data, a resolution factor of picking is computed considering the balanced inverse velocity standard deviation in the dispersion patterns according to the frequency, as described by O'Neill (2003).

Typical seismic shot gathers before and after dynamic compaction, with the related dispersion f - c images showing the fundamental mode, are presented in Fig. 5(a,b). One can see here the c for high frequencies (from 15 Hz) after dynamic compaction. The initial models for the inversion process of dispersion curves were obtained by converting f - c curves and the resolution factor as a function of the frequency in the depth-velocity domain ($depth$ - c) using the empirical relation of a half wavelength

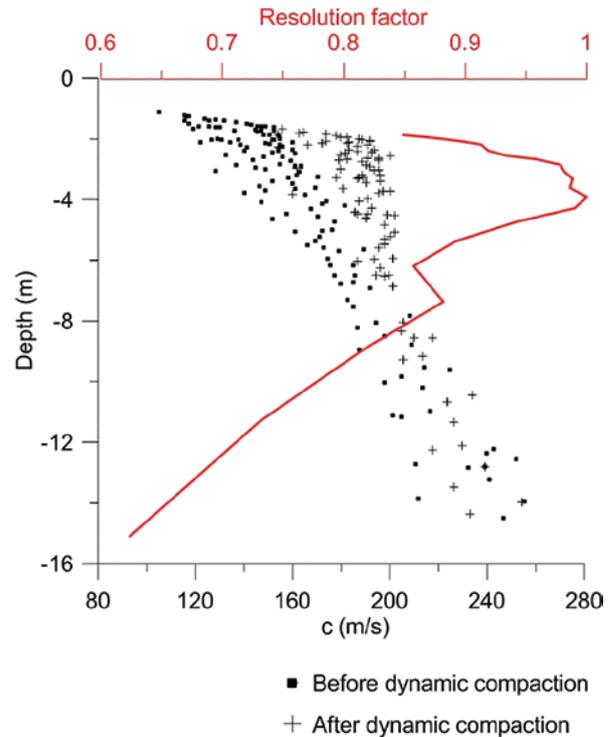


FIGURE 6

$Depth$ - c converted dispersion curves used as initial models before and after dynamic compaction obtained along the seismic lines A, B, C and D. The mean resolution factor of dispersion curves picking overlays the plot in red solid lines. We observe the increased c (around 40 m/s) from 0–8 m depth after dynamic compaction.

described by Xia *et al.* (1999a). Figure 6 shows the initial models obtained before and after dynamic compaction using this relation and the mean resolution factor of dispersion curves picking. On this figure we observe increased c (around 40 m/s) from 0–8 m depth after dynamic compaction. Below 8 m depth, c appears to be more or less unchanged. The mean resolution factor shows that the picking accuracy of the dispersion curves decreases with depth. According to this observation, a lower confidence is given to estimated velocities at high depths.

Each dispersion curve was then individually inverted into a 1D $depth/V_s$ profile. A 2D matrix of V_s was produced by gathering all 1D V_s profiles into sequential order according to the surface coordinate at the midpoint of the spread, the common procedure described by Park (2005). Ten iterations for inverting velocities and ten more for inverting layers thickness were necessary to converge to the final model. This allowed a root mean square error between the measured and calculated dispersion curves less than 5%. The LLS inversion technique used for this study is adapted from Hermann (1987).

CONE PENETRATION TEST

As described in Lunne *et al.* (1997), in the CPT, a cone on the end of a series of rods is pushed into the ground at a constant

rate and continuous or intermittent measurements are made of the resistance to penetration of the cone. Measurements are also made of either the combined resistance to penetration of the cone and outer surface of a sleeve or the resistance of a surface sleeve. The total force acting on the cone, Q_c , divided by the projected area of the con, A_c , produces the measured cone resistance, q_c .

Early applications of CPT mainly determined the soil geotechnical property of bearing capacity. The original cone penetrometers involved simple mechanical measurements of the total penetration resistance of pushing a tool with a conical tip into the soil. Different methods were employed to separate the total measured resistance into components generated by the conical tip and friction generated by the rod string. In the 1960s, a friction sleeve was added to quantify this component of the friction and aid in determining soil cohesive strength (Begemann 1965). Electronic measurements began in 1948 and improved further in the early 1970s (De Reister 1971). Most modern electronic CPT cones now also employ a pressure transducer with a filter to gather pore water pressure data. CPT testing, which also gathers these piezometer data, is called CPTU testing. CPT and CPTU testing equipment generally advances the cone using hydraulic rams mounted on either a heavily ballasted vehicle or using screwed-in anchors as a counter-force. One advantage of CPT over the Standard Penetration Test (SPT) is a more continuous profile of soil parameters, with CPTU data recorded typically at

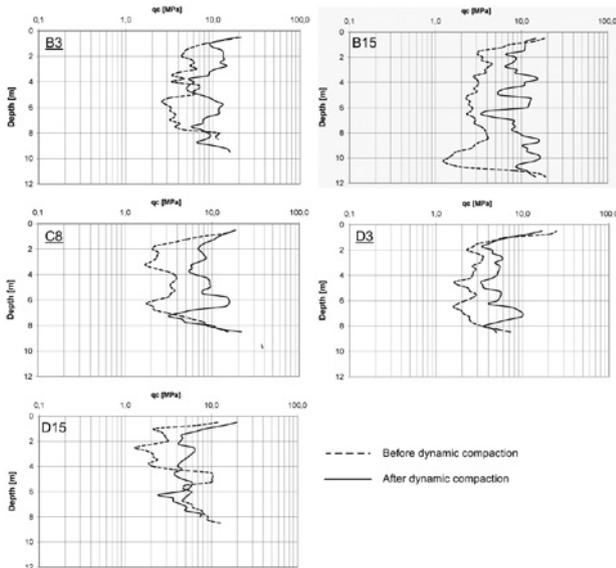


FIGURE 7 q_c soundings obtained at the points B3, B15, C8, D3 and D15 before and after dynamic compaction. Underlined items correspond to q_c soundings for which V_s profiles are shown in Fig. 8. Here we observe an increased q_c (2–9 MPa) for almost all soundings from 0 to 8–10 m depth after dynamic compaction. Let us note that points B3 and D15 highlight slices between 4–6 m depth with no difference between q_c before and after dynamic compaction.

2 cm intervals. CPT for geotechnical applications was standardized in 1986 by ASTM Standard D 3441 (ASTM 2004).

The q_c reference soundings before compaction were established in September 2009 at five points (B3, B15, C8, D3 and D15) on the test area (Fig. 2). The second CPTs were carried out at the end of October 2009, after heavy dynamic compaction at the same points in order to assess the effect of compaction. Figure 7 shows the q_c soundings obtained before and after dynamic compaction. On this figure we observe the increased q_c (2–9 MPa) for almost all soundings from 0 to 8–10 m depth after dynamic compaction. Below these depths, q_c seems to be unchanged.

DATA ANALYSIS

Inverted V_s profiles at the points B3, C8 and D15 before and after dynamic compaction are shown on Fig. 8. This figure indicates that V_s increased for about 10–80 m/s at a depth of 1–8 m. The

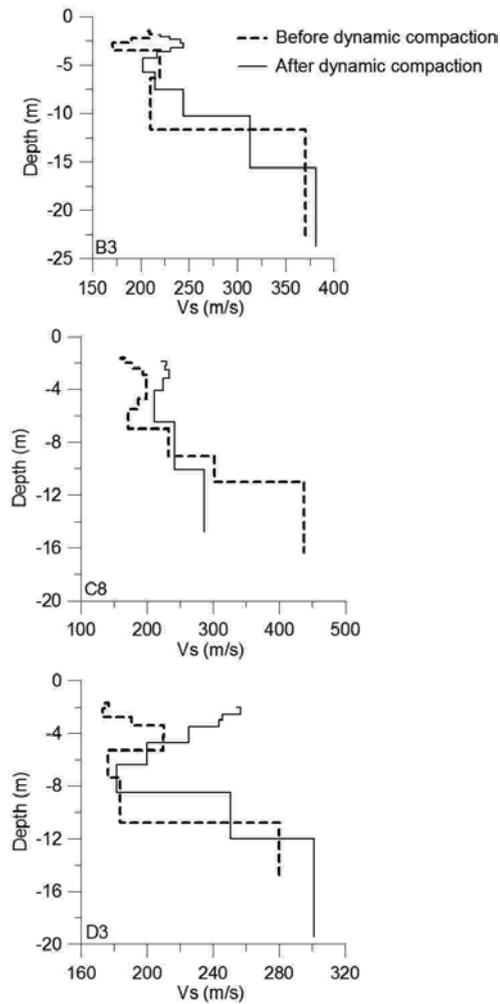


FIGURE 8 V_s profiles at the points B3, C8 and D15 before and after dynamic compaction. This figure indicates that V_s increased for about 10–40 m/s at a depth of 1–8 m.

B3 profile shows that V_s at the 4–6 m depth slice did not increase as already observed on the B3 q_c sounding. In general, we can observe consistency in the shape and maximum depth of increased q_c and V_s before and after dynamic compaction at the points B3, C8 and D15, which provides a consistency check of V_s sensibility to improved mechanical parameters in the near-surface. However V_s profiles show a global lack of consistency below 8 m depth. This can be explained by the low resolution of the picking of dispersion curves for low frequencies, which results in V_s inconsistencies in depth (Fig. 6). Following this observation, lower confidence is given to estimated V_s with depth.

In order to assess the quality of ground compaction, the variables dV_s and dq_c , corresponding respectively to the difference between inverted V_s and q_c after and before compaction, were considered. The relation dV_s-dq_c is plotted in Fig. 9.

For many engineering applications, it is common practice to normalize measured q_c and SPT N-values to a reference effective overburden stress, typically 1 atmosphere (approximately 101 kPa) (Robertson and Wride 1998; Karray *et al.* 2010; Harutoonian *et al.* 2012). Several studies have concluded that use of stress-normalized N - or q_c - values in V_s correlations prove to be considerably less accurate than correlations based on non-normalized values (Sykora and Stokoe 1983; Lodge 1994; Piratheepan 2002; Hasancebi and Ulusay 2007; Wair *et al.* 2012). Additionally, for the purpose of site classification in accordance with design codes and calculation of V_{s30} , it is not appropriate to normalize penetration resistance for overburden stress. For some applications, such as liquefaction triggering assessment, it may be necessary to normalize V_s estimates to a reference stress level. In such cases, V_s can be estimated from non-normalized penetration resistance and then normalized for overburden.

The representative q_c was determined as an average value of q_c between depths corresponding to the V_s layers for profiles B3, C8 and D15 where both V_s and q_c values were available. This

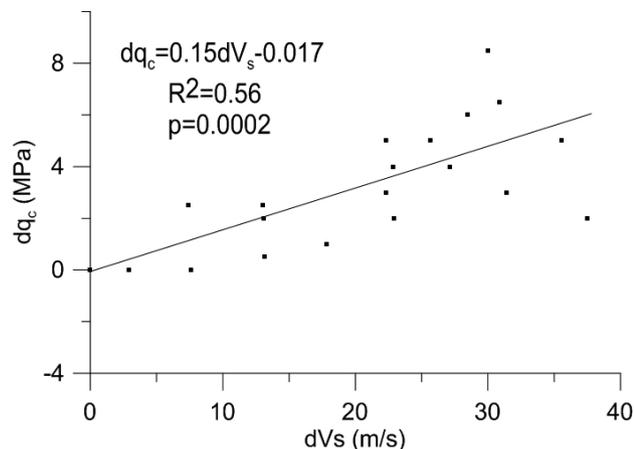


FIGURE 9 Relation dV_s-dq_c . V_s increases as the representative q_c increases, showing an acceptable correlation among variables ($R^2=0.56$).

procedure allowed the extraction of 20 samples used in a linear regression between dV_s and dq_c . dV_s increases as the representative dq_c increases, showing an acceptable correlation among variables ($R^2=0.56$). This indicates that the shear-wave velocity determined by the MASW test is a meaningful variable in the evaluation of the quality of layer compaction. Figure 9 shows the linear regression between dq_c and dV_s from which dq_c can be obtained using the following equation:

$$dq_c = 0.15dV_s - 0.0173 \quad (2)$$

with a p value of 0.0002 indicating that the regression is statistically significant. Calculating the spatial variation of dq_c as a function of dV_s values using equation (2) then becomes possible over entire seismic lines.

RESULTS

dq_c sections obtained using equation (2) along the lines A, B, C and D are shown in Fig. 10. The results indicate that the near-surface, globally less than 5–8 m depth, was homogeneously improved by dynamic compaction according to a q_c increase of 4–7 MPa. We observe higher compaction depths (around 10–13 m) between the abscissa 80–120 m on the lines A, B and C that may be interpreted as former channels of the Gier river. Let us note that the maximum compaction depths with the CPTs are consistent with these observed limits. Below, materials were not improved by compaction. This can be explained by the presence of compact sand and gravel constituting the geotechnical bedrock. However, on lines B, C and D, an improved layer (1–4 MPa) is visible from 10–15 m depth in the west part of the test area probably due to the presence of intercalated soft sand.

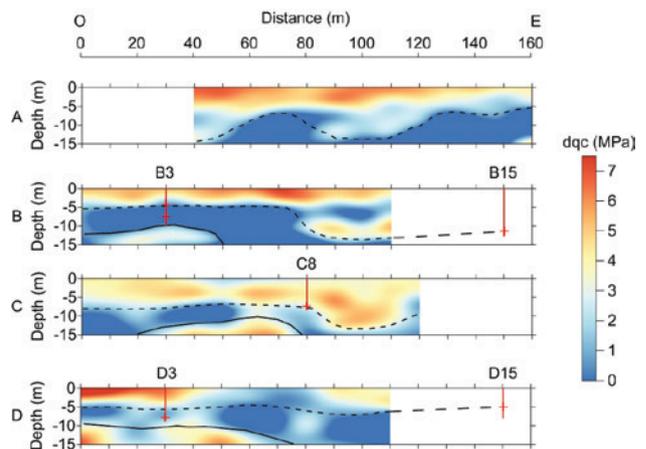


FIGURE 10 dq_c sections obtained using equation (2) along the lines A, B, C and D. Maximum compaction depths observed on CPTs are represented using red crosses. The improved near-surface appears in shades of red. The interpreted maximum compaction depths using the 3 MPa threshold overlay the sections in black dotted lines. Black solid lines represent buried improved zones due to intercalated soft sand.

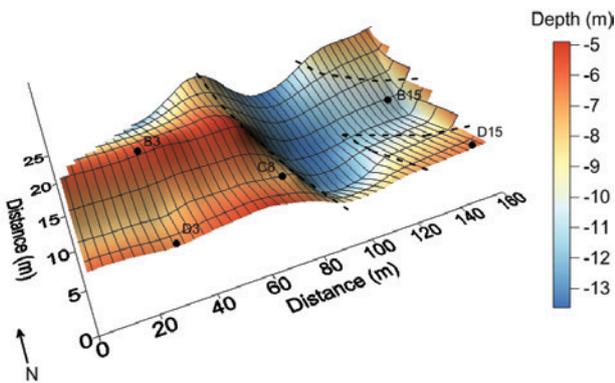


FIGURE 11

3D view of the interpreted maximum ground compaction depth over the entire test area. Potential routes of a former channel of the Gier river overlay the surface in black dotted lines.

According to these observations, an interpreted maximum compaction depth, assimilated to the top of the bedrock, was extracted using 3 MPa thresholds on dq_c sections and compaction depths from CPTs, highlighted in Fig. 10. In order to assess its variability over the entire test area, interpreted maximum ground compaction depths were then interpolated using the kriging technique (Fig. 11).

CONCLUSION

An evaluation procedure to determine the quality and extent of ground compaction is proposed by effectively combining the results from *in situ* MASW tests with those from CPT tests. The following conclusions can be drawn: (i) the MASW method, which is rapid and non-intrusive, is an effective tool in assessing the quality and depth of ground improvement by measuring V_s profiles before and after compaction, (ii) comparison between the MASW and CPTs results, before and after compaction, demonstrates the correlation between dV_s and dq_c (iii) a procedure for estimating the improvement of the ground due to compaction is proposed by combining *in situ* V_s profiles determined by MASW tests with the linear regression between dV_s and dq_c determined by CPT tests, (iv) results demonstrate the ability of the procedure to accurately delineate near-surface features and detect bedrock. This study demonstrates the potential of surface waves in geotechnical investigations.

ACKNOWLEDGEMENTS

This work was carried out under BRGM financial contribution. The authors thank Ménard France for permission to publish the results. We also thank the Colorado School of Mines for allowing the SU package at geophysicists' disposal in freeware (Cohen and Stockwell 1997) and R. Herrmann for placing the Computer Program in Seismology free on the web. The authors thank C.B. Park, D. Albarello and an anonymous reviewer for their constructive and detailed reviews.

REFERENCES

- Addo K.O. and Robertson P.K. 1992. Shear-wave velocity measurements of soils using Rayleigh waves. *Canadian Geotechnical Journal* **29**, 558–568.
- ASTM. 2004. Standard Method of Deep Quasi-Static Cone and Friction-Cone Penetration Tests of Soil. ASTM Standard D 3441. ASTM International, West Conshohocken, PA, 7.
- Begemann H.K.S. 1965. The Friction Jacket Cone as an Aid in Determining the Soil Profile. *Proceedings of the 6th ICSMFE*, Montreal, Quebec, Canada, Vol I, 17–20.
- Bitri A., Le Bégat S. and Baltassat J.M. 1998. Shear wave velocity determination of soils from in-situ Rayleigh waves measurements. *Proceedings of the 4th Meeting EEGS*, Barcelona, Spain, 503–506.
- Brûlé S., Javelaud E.H., Ohmachi T., Nakamura Y. and Inoue S. 2010b. H/V method used to qualify the modification of dynamic soil characteristics due to ground improvement work by means of heavy compaction process. A case study: the former Givors's glass factory area. *7th International Conference on Urban Earthquake Engineering and 5th International Conference on Earthquake Engineering*, Tokyo, Japan, 02–26, 451–455.
- Cohen J.K. and Stockwell Jr. J.W. 1997. CWP/SU: Seismic Unix. A free package for seismic research and processing. Center for Wave Phenomena, Colorado School of Mines, Golden, CO.
- De Reister J. 1971. Electric Penetrometer for Site Investigations. *Journal of SMFE Division, ASCE* **97**(SM-2), 457–472.
- Dorman J. and Ewing M. 1962. Numerical inversion of seismic surface waves dispersion data and crust-mantle structures in the New-York-Pennsylvania area. *Journal of Geophysical Research* **67**, 5227–5241.
- Forbriger T. 2003a. Inversion of shallow seismic wavefield: I. Wavefield transform. *Geophysical Journal International* **153**, 719–734.
- Forbriger T. 2003b. Inversion of shallow seismic wavefield: II. Inferring subsurface properties from wavefield transforms. *Geophysical Journal International* **153**, 735–752.
- Gabriels P., Sneider R. and Nolet G. 1987. In-situ measurements of shear wave velocity in sediments with higher mode Rayleigh waves. *Geophysical Prospecting* **35**, 187–196.
- Harutoonian P., Leo C.J., Doanh T., Castellaro S., Zou J.J., Liyanapathirana D.S. *et al.* 2012. Microtremor measurements of rolling compacted ground. *Soil Dynamics and Earthquake Engineering* **41**, 23–31.
- Hasancebi N. and Ulusay R. 2007. Empirical correlations between shear wave velocity and penetration resistance for ground shaking assessment. *Bulletin of Engineering Geology and the Environment* **66**, 203–213.
- Hermann R.B. 1987. *Computer Programs in Seismology*. Saint-Louis University, USA.
- Karray M., Lefebvre G., Ethier Y. and Bigras A. 2010. Assessment of deep compaction of the Péribonka dam foundation using 'modal analysis of surface waves' (MASW). *Canadian Geotechnical Journal* **47**, 312–326.
- Kim D.S. and Park. H.C. 1999. Evaluation of ground densification using spectral analysis of surface waves (SASW) and resonant column (RC) tests. *Canadian Geotechnical Journal* **36**, 291–299.
- Lodge A.L. 1994. *Shear wave velocity measurements for subsurface characterization*. PhD thesis, Dept of Civil and Environmental Engineering, University of California, Berkeley, CA.
- Lunne T., Robertson P.K. and Powell J.J.M. 1997. *Cone Penetration Testing in Geotechnical Practice*. Blackie Academic/Routledge Publishing, New York.
- Luo Y., Xia J., Miller R.D., Xu Y., Liu J. and Liu Q. 2008. Rayleigh-wave dispersive energy imaging using a high resolution linear Radon transform. *Pure and Applied Geophysics* **165**, 903–922.
- McMechan G.A. and Yeldin M.J. 1981. Analysis of dispersive waves by wave field transformation. *Geophysics* **49**, 1169–1179.

- Moktar T.A., Hermann R.B. and Russel D.R. 1988. Seismic velocity and Q model for the shallow structure of the Arabian shield from short-period Rayleigh waves. *Geophysics* **53**, 1379–1387.
- Nazarian S. and Stokoe K.H. 1984. In-situ shear wave velocities from spectral analysis of surface waves. *Proceedings 8th World Conference on Earthquake Engineering* **3**, 31–38.
- O'Neill, A. 2003. *Full-waveform reflectivity for modeling, inversion and appraisal of seismic surface wave dispersion in shallow site investigations*. MSc thesis, The University of Western Australia.
- Park C.B. 2005. MASW – Horizontal resolution in 2D shear-velocity (V_s) mapping. KGS Open-File Report 2005–4.
- Park C.B., Miller R.D. and Xia J. 1998. Imaging dispersion curves of surface waves on multi-channel record. Technical Program with Biographies. *SEG 68th Annual Meeting*, New Orleans, LA., 1377–1380.
- Park C.B., Miller R.D. and Xia J. 1999. Multi-channel analysis of surface waves. *Geophysics* **64**(3), 800–808.
- Park C.B., Miller R.D., Xia J. and Ivanov J. 2007. Multi-channel analysis of surface waves. (MASW) active and passive methods. *Leading Edge* **26**, 60–64.
- Piratheepan P. 2002. *Estimating shear wave velocity from SPT and CPT data*. MSc Thesis, Clemson University.
- Richart Jr.F.E., Hall Jr.J.R. and Woods R.D. 1970. *Vibration of Soils and Foundations*. Prentice Hall Inc., Englewood Cliffs, N.J.
- Robertson P.K. and Wride C.E. 1998. Evaluating cyclic liquefaction potential using the cone penetration test. *Canadian Geotechnical Journal* **35**, 442–459.
- Romdhan A., Grandjean G., Brossier R., Rejiba F., Opero S. and Virieux J. 2011. Shallow-structure characterization by 2D elastic full-waveform inversion. *Geophysics* **76**, (R81). doi:10.1190/1.3569798
- Song Y.Y., Castagna J.P., Black R.A. and Knapp R.W. 1989. Sensitivity of near-surface shear-wave velocity determination from Rayleigh and Love waves. Technical program with Biographies. *SEG 59th Annual Meeting*, Dallas, TX, 509–512.
- Sykora D.E. and Stokoe K.H. 1983. Correlations of in-situ measurements in sands of shear wave velocity. *Soil Dynamics and Earthquake Engineering* **20**, 125–136
- Wair B.R., De Jong J.T. and Shantz T. 2012. Guidelines for estimation of shear wave velocity profiles. Pacific Earthquake Engineering Research Center, PEER 2012/08.
- Xia J., Miller R.D. and Park C.B. 1999a. Estimation of near-surface shear-wave velocity by inversion of Rayleigh waves. *Geophysics* **64**(3), 691–700.
- Xia J., Miller R.D. and Park C.B. 1999b. Advantage of calculating shear-wave velocity from surface waves with higher modes. *SEG 70th Meeting*, Houston, Texas.



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