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1 **Dynamic rupture of cross-cutting faults: A possible rupture process of the**

2 **2007 Mw6.6 Niigata-Ken Chuetsu-Oki earthquake**

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**Abstract**

**The 2007 Mw6.6 Niigata-Ken Chuetsu-Oki earthquake implies a complex fault mechanism. Several surveys suggest the possibility of two main segments crosscutting each other at the middle, namely north-western dipping in the north and south-eastern dipping in the south. We numerically model dynamic rupture propagation along the inferred segmented fault system using a boundary integral equation method (BIEM). The possibility of the rupture transfer is numerically shown and two rupture modes are reported. Simultaneous rupture transfer along the overlapping part is possible only if stress is highly loaded, however this rupture mode provides a too much slip amount in the result. In the other case where regional stress is relatively low but pore pressure is high enough to govern the rupture criterion (described as low frictional coefficient case in this study), the rupture transfer to the other segment does not occur until the rupture terminates on the first segment regardless of the crosscutting distance between two segments. In this case at the middle cross-cutting part, a possible rupture scenario is that there occur rupturings sequentially in different directions; first on the north-western dipping segment by southbound rupture and on the south-eastern dipping segment by northbound rupture a few seconds later. The simulated rupture scenario on conjugated faults can be strongly led by pre-existing fault structure formed by both the Miocene rifting during an opening stage of the Japan Sea and the subsequent shortening of the crust. Our results infer the importance of investigating earthquake rupture scenarios in complex fault system based on the geophysical and geological information, and the dynamic rupture mechanics.**

# 1. Introduction

How does material rupture under a load? For figuring out the problem clearly, let us think a two-dimensional profile where there are two potential orientations of fault plane, called “conjugate”, with respect to the loaded principal force (e.g. Scholz, 2002). Usually all materials break selectively along either dominant plane but they do not rupture on both planes simultaneously. Once one plane becomes dominant in the material, the other plane is no longer activated in terms of stress concentration.

Earthquakes are one of the largest rupture phenomena in a solid material. Fault system has already evolved intensively in the Earth’s interior, by interacting and evolving as a geodynamical system over different scales of time and space, so that the characteristics of the materials and the rupture mechanism of earthquakes can be more complex than observed in rock specimens in laboratory experiments. Geologists have pointed out the role played by fault geometry on earthquake rupture (e.g. King and Nabelek, 1985). This can be quantitatively simulated as a dynamic rupture process, as shown in the case of the Mw7.2 1992 Landers, California earthquake by Aochi and Fukuyama (2002), for example. Although this earthquake shows a complex geometry, with bends, branches and segmentations, it can be treated as a single fault plane from the viewpoint of standard seismological description (Olsen et al., 1997).

The Mw6.6 Niigata-ken Chuetsu-Oki earthquake occurred on 16 July 2007 in north-eastern Japan (Figure 1). The hypocenter was located under the Sea of Japan at a depth of around 12 km and several kilometres from the coastline of the main island. The ruptured area expanded

1 over a length of about 30 km, according to various studies, and part of the rupture may have  
2 occurred beneath the main island. Although the focal mechanism clearly showed reverse  
3 faulting, it was difficult to determine the fault plane only from the seismological waveform  
4 inversion of finite source parameters (Koketsu et al., 2007; Aoi et al., 2008). Extensive  
5 seismological and geophysical surveys have been carried out to record the aftershocks for  
6 imaging the event distribution and geophysical structure, and to construct the source model.  
7 The final results determined by the dense temporal seismic network deployed after the  
8 mainshock show a clear variation in aftershock distributions on cross-sections from north to  
9 south along the fault strike, as shown in Figure 1 (Kato et al., 2008; Shinohara et al., 2008),  
10 with a north-western dip in the north but a south-eastern dip in the south. Furthermore, fine  
11 tomography studies elucidate that there is a segment boundary zone situated between the  
12 north and the south, which is coincident with the complex aftershock zone where many  
13 conjugate fault planes exist (Kato et al., 2008). Although a southeastward-tilted structure of  
14 the basement is dominant in the south, a northwestward-tilted structure of the basement  
15 gradually develops toward the north (Kato et al., 2008, 2009). It is considered that the lateral  
16 complex variations in basement structure along the fault strike have been formed by both the  
17 Miocene rifting during the opening stage of the Japan Sea and the subsequent shortening of  
18 the crust (e.g., Okamura et al., 1995; Kato et al., 2009). These results indicate that two  
19 ruptured segments could be cross-cutting each other in the middle (CT-section in Figure 1).  
20 The analyses of geodetic deformation based on InSAR images also support the segmented  
21 fault models, mainly because of the irregular pattern of surface deformation from north to  
22 south (Aoki et al., 2008; Nishimura et al., 2008, 2009). Recently, Takenaka et al (2009) has  
23 demonstrated that the mainshock rupture initiated and propagated on the northwest-dipping  
24 plane to the updipward, using the initial rupture phase of the P-wave portion. This

1 seismological evidence suggests the sequential or simultaneous rupture of the two conjugate  
2 fault planes.

3 From a geological point of view, fault traces can cross each other so as to be consistent with  
4 orientations of the regional tectonic stress, although this does not assure that earthquake  
5 rupture can occur simultaneously on both fault traces. For example, the 1994 M6.8 Northridge  
6 and the 1971 M6.6 San Fernando earthquakes occurred on adjacent thrust faults, which are  
7 dipping in opposite directions, i.e. conjugate (Carena and Suppe, 2002). However, there is  
8 little evidence of simultaneous activation of two conjugate fault planes. The 1997 Kagoshima,  
9 Japan, earthquake (M6.0) shows the possibility of simultaneous rupture on conjugated,  
10 adjacent strike-slip segments (Horikawa, 2001). This configuration has been numerically  
11 simulated as two separated segments (Kase and Kuge, 2001). In the case of the 2000 Tottori,  
12 Japan, earthquake (M6.7), some secondary segments perpendicular to the main traces  
13 segments were identified and numerically modelled (Fukuyama et al., 2003; Fukuyama, 2003).  
14 However, they remain secondary. Thus the 2007 Niigata-ken Chuetsu-Oki earthquake is a  
15 unique occasion to observe simultaneous or sequential rupture on two-conjugated fault planes  
16 crosscutting in the middle. This study aims to understand how this is possible from a  
17 mechanical viewpoint, through numerical simulations.

18

## 19 **2. Numerical Method and Fault Models**

20

21 We adopt in this study a boundary integral equation method (BIEM) (Aochi et al., 2000) for  
22 simulating the dynamic rupture process on complex fault geometry with a given rupture  
23 criterion. This method is powerful for various elasto-dynamic problems, especially including  
24 discontinuous cracks, which can spontaneously propagate along a given crack geometry, or in

1 the case of an unknown geometry (e.g. Kame and Yamashita, 1999). As the mainshock  
2 hypocenter depth is around 12 km on the north-western dipping fault plane (Figure 1) and  
3 most of aftershocks are distributed within the basement in a depth range of 6-15 km where the  
4 P-wave velocity is greater than 6 km/s (e.g., Kato et al., 2008; Shinohara et al., 2008), it is  
5 sufficient to model this earthquake in an infinite medium without taking into account of the  
6 ground surface effect. As the method is strictly based on the theory of infinitesimal  
7 deformation in linear elasticity, no numerical nuisance occurs at crack junctions, as  
8 successfully shown for crack branching (Tada and Yamashita, 1997; Aochi et al., 2002). In  
9 the discretized formulation of the BIEM, stress is always evaluated at the center of each  
10 planar discretized element where slip velocity is assumed constant during a time step. We are  
11 able to generate the fault geometry model avoiding the kinematic problem of at any jointing  
12 points between elements. Such configurations are still difficult to study with other volumetric  
13 numerical methods, because a junction point should have some values of displacement, which  
14 is not always consistent with the kinematics (e.g. Andrews, 1989).

15  
16 Figure 2 shows the fault model to be simulated in this study. The northern segment (Segment  
17 1) is 10.4 km long and 8 km wide and the southern segment (Segment 2) is 20 km long and  
18 14.4 km wide. We calibrate these geometries based on Figure 1, namely the relocated  
19 aftershock distribution and the hypocenter of the mainshock (Kato et al., 2008). Nevertheless,  
20 there remains slight uncertainty especially in the dip angle ( $\delta$ ) of Segment 2 and the  
21 crosscutting distance ( $L$ ) of the two segments. The former is due to a lower resolution for off-  
22 shore analysis and the latter is too small to be determined seismologically. The latter is  
23 because it is always difficult to determine precisely the fault extension only from the  
24 aftershock distribution. Therefore we vary these two parameters in the simulations. The other  
25 model parameters are summarized in Table 1. The medium property is consistent with the

1 tomographic image of velocity near the earthquake source (Kato et al., 2008). The simulation  
2 is carried out on a parallel computer, requiring only about 10 minutes for the first 200 time  
3 steps using 8 cores.

4  
5

## 6 **3. Rupture criterion and stress field**

### 7 ***3.1 Slip-weakening law and Coulomb law***

8 The analysis of focal mechanisms of the aftershocks implies that the maximum principal  
9 stress axis is on the horizontal plane and perpendicular to the fault strike (Kato et al.,  
10 unpublished data, 2008; systematical analysis reported on the NIED homepage,  
11 <http://fnet.bosai.go.jp>, 2007). It is consistent with the focal mechanism of the main shock and  
12 regional tectonic setting (Kato et al., 2006; Townend and Zoback, 2006). Therefore, we  
13 assume that the fault slip direction is purely along the dip, which is also indicated by the focal  
14 mechanism of the main shock (Figure 1). On the other hand, this indicates that the main shock  
15 releases a part of the accumulated regional stress and does not rotate the stress field. The  
16 rupture process is governed by a constitutive relation between on-going fault discontinuity  
17  $\Delta u$  and shear stress  $\tau$  called as slip-weakening law, which is widely accepted in dynamic  
18 rupture process of earthquakes (e.g. Ida, 1972, Palmer and Rice, 1973). We adopt the  
19 following standard mathematical expression:

20

$$21 \quad |\tau(\Delta u)| = \tau_r + (\tau_p - \tau_r)(1 - |\Delta u|/D_c)H(1 - |\Delta u|/D_c) \quad (1)$$

22

1 where  $H(\dots)$  is a Heaviside function,  $\tau_p$  and  $\tau_r$  are peak and residual stresses and  $D_c$  is  
2 critical slip-weakening displacement, also shown in Figure 3. Here we explicitly write down  
3 Equation (1) with absolute values, because loaded shear stress on two segments has  
4 mechanically opposite sign so that slip direction is reversed in a mechanical view. This  
5 corresponds to the seismological description where fault strike is in reverse although fault  
6 rake is the same. Before the onset of dynamic rupture, namely in the case of  $|\Delta u| = 0$ , shear  
7 stress  $|\tau|$  remains lower than  $\tau_p$ .  $\tau_p$  and  $\tau_r$  are given by the Coulomb law, namely a product  
8 of the frictional coefficient ( $\mu_s$ : static or  $\mu_d$ : dynamic frictional coefficients) and applied  
9 normal stress  $\sigma_n$  as:

10

$$11 \quad \tau_p = \mu_s \times \sigma_n, \text{ and} \quad (2)$$

$$12 \quad \tau_r = \mu_d \times \sigma_n \quad . \quad (3)$$

13

14 In the following simulations, fault strength and applied stress is determined uniquely by the  
15 assumed tectonic regional stress assuming a single value of  $\mu_s$  in a simulation without  
16 privileging any direction of fault geometry. As  $\tau_p$  and  $\tau_r$  change significantly according to  
17 the assumption, we give the same slip-weakening rate ( $\equiv (\tau_p - \tau_r) / D_c$ ) for all the cases  
18 instead of fixing a value of  $D_c$ , namely = 30 [MPa/m], which is consistent with the value  
19 inferred from the seismological analyses (e.g. Ide and Takeo, 1997; Olsen et al., 1997). This  
20 also allows us to initiate the dynamic rupture in the same way, as the rupture propagation  
21 criteria is controlled by the slip-weakening rate (Matsu'ura et al., 1992; Madariaga and Olsen,  
22 2000). We give Segment 1 an initial circular crack of a radius of 1 km where the strength is

1 supposed to drop instantaneously at the beginning of the simulation in order to initiate the  
2 spontaneous dynamic rupture propagation.

3

### 4 **3.2 Tri-axial compression stress field**

5 We assume a uniform compressional tectonic stress around the source area. The maximum  
6 principal stress is taken to be horizontal, perpendicular to the fault strike, and the minimum  
7 one is vertical, namely pressure by gravity. The intermediate one should be horizontal,  
8 parallel to the fault strike, but as we suppose a pure reverse faulting, this component does not  
9 play a role in the dynamic modelling. The initial shear and normal stresses applied on the  
10 fault segments are given respectively by:

11

$$12 \quad \tau_0 = \frac{\sigma_1 - \sigma_3}{2} \sin 2\Phi, \text{ and} \quad (4)$$

$$13 \quad \sigma_n^0 = \frac{\sigma_1 + \sigma_3}{2} - \frac{\sigma_1 - \sigma_3}{2} \cos 2\Phi, \quad (5)$$

14

15 letting  $\sigma_1$  and  $\sigma_3$  be the maximum and minimum principal stresses (compression taken as  
16 positive), and  $\Phi$  be the angle of the fault segment inclined with respect to the  $\sigma_1$  axis. The  
17 model parameters are summarized in Table 1. In the mechanical coordinate we take, Equation  
18 (4) gives a positive shear stress loading on Segment 1, while it gives a negative loading on  
19 Segment 2. This means that, in the same coordinate, fault slip is positive on Segment 1 but  
20 negative on Segment 2. Hereafter the simulations results are presented based on this sign  
21 convention.

22

1 What is known is the minimum principal stress  $\sigma_3$  in the vertical direction. We take the  
 2 hydrostatic condition at a depth of  $z = 8 \text{ km}$  where cross-cutting section is located. We precise  
 3 that all the normal stresses are effective normal stresses, since the effect of pore pressure is  
 4 subtracted. Therefore,

$$5 \quad \sigma_3 = P(z = 8 \text{ km}) = (\rho - \rho_H)gz = 117.6 \text{ MPa}, \quad (6)$$

8 where  $\rho$  is the material density (Table 1),  $\rho_H$  is the water density and  $g$  is the gravity  
 9 constant. We suppose  $\mu_s = 0.7$  as inferred from the laboratory experiments (e.g. Byerlee,  
 10 1978) and illustrate the Mohr circle so as not to exceed the rupture criterion imposed by the  
 11 frictional line defined by  $\mu_s = 0.7$  (Figure 4(a)). The angle most favourable to the rupture is  
 12  $\Phi = 27.5^\circ$  from simple algebra. In this case as summarized in Table 2, we obtain  $\sigma_1 = 433.9$   
 13 MPa and a strength excess on Segment 1 of 56.5 MPa. In Figure 4, for reference, we add  
 14 dynamic friction (residual stress) lines of  $\mu_d = 0.33$  which gives the S value of 1 on Segment  
 15 1, defined by

$$16 \quad S = \frac{\tau_p - |\tau_0|}{|\tau_0| - \tau_r}, \quad (7)$$

19 (Das and Aki, 1977), namely strength excess is equal to static stress drop (at the initial normal  
 20 stress). This condition is quite natural for allowing the dynamic rupture on a plane fault in the  
 21 term of the value S ( $S \approx 1$ ), but it should be noted that  $\mu_d$  is half of  $\mu_s$ . In other words, about  
 22 half of the absolute stress is relieved due to the earthquake rupture and its amount is an order  
 23 of 50 MPa, which seems to be still high comparing to the averaged value inferred from the

1 natural earthquakes. If we taken a higher value of  $\mu_d$ , for example,  $\mu_d = 0.7\mu_s$ , S value is 7.9  
2 for Segment 1, which does not allow the initiation of rupture in our situation.  
3  
4 The fault strength may be much weaker for example due to over-pressure of existing fluid (e.g.  
5 Miller et al., 2004; Sibson, 2007). Actually the fluid existence is inferred at the lower crust  
6 beneath the fault system from the tomographic interpretation (Kato et al., 2009). Although we  
7 do not aim to discuss the detail of such complex mechanism, we instead present it by lower  
8 frictional coefficients. Figure 4(b) shows the stress configuration assuming  $\mu_s = 0.3$ . The  
9 most favourite angle of fault plane is  $\Phi = 36.7^\circ$ . In this case, we get  $\sigma_1 = 212$  MPa and  $\mu_d =$   
10  $0.24 \approx 0.6 \mu_s$  for the condition of  $S = 1$  on Segment 1.

11  
12 It is worth of discussing different parameter settings in a fault system of complex geometry as  
13 the effect of fluid overpressure is not yet quantified in the earthquake rupture. We test three  
14 cases,  $\mu_s = 0.7, 0.5$  and  $0.3$ . As explained above (e.g. Figure 4) and summarized in Table 2, a  
15 higher value of the frictional coefficient corresponds to a high stress condition (strong fault  
16 strength), while a lower value implies a lower stress (strength) condition with significant pore  
17 fluid effect. We have not introduced here any depth-dependence on the model parameters.  
18 However the fault system is located at depth enough in the seismogenic zone so as not to be  
19 suffered from the significant depth-dependency in the shallow or subsurface crustal structures.  
20 We recall that the main objective of this paper is to investigate the mechanics of rupture  
21 transfer from one segment to another along the intersection which is located at the same  
22 seismogenic depth from north to south. This is why our parameter choice does not lose the  
23 generality of mechanics.

24

## 4. Simulation results

Is it possible that the crosscutting conjugate faults rupture simultaneously? If so, how does it occur? For investigating the possible mechanism, we make three parameters variable as explained in the previous sections, namely the dip ( $\delta$ ) of Segment 2, the overlapping length ( $L$ ) of two segments and the static frictional coefficient ( $\mu_s$ ). We carry out 45 simulations in total. Figure 5 shows snapshots of two typical dynamic rupture scenarios found among the numerical simulations, and Figure 6 summarises the scenarios by parameters.

The first example (case  $\mu_s = 0.7$ ,  $\delta = 35^\circ$  and  $L = 4$  km) shows that rupture transfer on Segment 2 occurs simultaneously during the rupture propagation on Segment 1 (Figure 5a). The triggering is usually caused by strong perturbation of dynamic stress field brought from Segment 1, and this scenario is due to the very low value of parameter  $S$  on Segment 2 (close to 0) as inferred from Figure 4. We found that the initial condition is much more favour on Segment 2 than Segment 1. This occurs commonly for the fault geometry of Segment 2 with a dip angle  $\delta$  of 30-40° indifferently from the frictional parameter  $\mu_s$  and the crosscutting distance  $L$  (except for  $L = 0$ ). The rupture initiates on Segment 2 immediately when a favourite stress wave (dynamically there are more or less both positive and negative changes) arrives and begin to propagate spontaneously. This kind of scenario is often found in our simulations of this study. This is due to the extreme condition naturally generated for the conjugate faults under a uniform stress field. In the case of parallel faults, for example, applied stress is the same for both so that it is difficult to favour Segment 2 (e.g. Harris and Day, 1999; Kase and Kuge, 2001). In the branched fault cases, it is reported that rupture can transfer to the secondary plane other than the principal plane where the rupture starts (Aochi

1 et al., 2002; Kame et al., 2003), however rupture mode is continuously the same in these cases  
2 (slip direction with respect to the fault direction). The particularity of our study is that slip  
3 direction is different on two segments, although they are both reverse faulting with respect to  
4 the geological description.  
5  
6 Another example (case  $\mu_s = 0.3$ ,  $\delta = 45^\circ$  and  $L = 4$  km) of Figure 5(b) presents sequential  
7 rupture propagation from Segment 1 to Segment 2. It is natural that this occurs when two  
8 segments are not overlapping each other in Figure 6 ( $L = 0$ ). More generally, this is observed  
9 when frictional coefficient  $\mu_s$  (namely stress) is lower (see again Figure 6), namely when  
10 stress concentration is required more to initiate rupture process on Segment 2. In this case, we  
11 emphasis on the fact that rupture on Segment 2 is not initiated during the rupture propagating  
12 along Segment 1 so that the rupture transfer is independent of the overlapping length of the  
13 segments. Figure 7 confirms the process in detail for the observed cases, showing the rupture  
14 time propagation on Segment 2. Segment 2 is not activated during the rupture propagation on  
15 Segment 1. It is called as shadow effect, or simply stress concentration is the most beyond the  
16 rupture front on the same plane. The rupture on Segment 2 begins from the point where the  
17 Segment 1 stops. It is then observed that the rupture on Segment 2 continues propagating to  
18 the south, while a part of the rupture comes back in the north direction on Segment 2 so as to  
19 fill the overlapping portion. In these cases, at the middle of this fault system where two faults  
20 planes are cross-cutting, two fault planes are ruptured with a time lag of a few seconds during  
21 a single earthquake. Namely, the north-western dipping rupture may have taken place first  
22 towards the south, and then the south-eastern dipping rupture starts later and propagates to  
23 both the north and the south.  
24

1 There is another scenario that rupture does not transfer on Segment 2 (Figure 5c). This is  
2 found when  $\delta = 50^\circ$ , the same dip as Segment 1, namely frictional and initial stress conditions  
3 are the same on both segments. As a result, it is in any case required that Segment 2 has a  
4 more favourable condition than Segment 1 in order for rupture to transfer.

5

## 6 **5. Discussion**

7 In the previous section, we have focused on the mechanical scenario of rupture transfer  
8 between conjugate segments. It is still difficult to argument which parameter setting adapts  
9 better with the 2007 Niigata-ken Chuetsu-Oki earthquake. Concerning about frictional  
10 coefficient  $\mu_s$ , any value allows rupture transfer on Segment 2. However it leads  
11 quantitatively different slip amount according to the allowed stress drop. In this meaning, a  
12 high  $\mu_s$  gives a too much slip more than 10 m with respect to the fault dimension. Note that  
13 rupture never begins at the hypocenter when a relatively high  $\mu_d$  is given (large S value).  
14 Therefore a low value of  $\mu_s$  adjust better the amount of slip expected for such fault dimension.  
15 About the fault geometry of Segment 2, a low dipping angle around  $30^\circ$  is inferred from the  
16 aftershock distribution (Kato et al., 2008; Shinohara et al., 2008), but they also note that this  
17 dip angle  $\delta$  may be more steep because of the incertitude of their hypocenter localization off  
18 shore. This is why we have varied this parameter. As long as the friction coefficient  $\mu_s$  is  
19 high, this angle  $\delta$  does not play a significant role as stress is loaded too much on Segment 2 of  
20 a wide range of dip angle  $\delta$ . For a smaller value of  $\mu_s$ , the possibility of rupture transfer does  
21 not change (a wide angle of  $\delta$  between  $30-45^\circ$ ). This is acceptable with respect to the  
22 incertitude of the observations. By the way, we do not have any direct indication on the  
23 crosscutting distance  $L$ , although the aftershock distribution (Figure 1) supports the cross-

1 cutting two segments at the middle. Our simulation shows the possibility of simultaneous  
2 rupture transfer for a high frictional coefficient, while rupture transfers sequentially one from  
3 one in other cases, for which it is essential that Segment 1 is terminated. Interpretation from  
4 the structural geology implies rather such discontinuity in fault geometry between the north  
5 and the south (Kato et al., 2008; 2009).

6

7 In this study, we have assumed a simple, but extreme Mohr–Coulomb criterion, in which the  
8 static friction line is tangential to the Mohr circle. Namely the applied stress reaches the  
9 rupture criterion at a direction ( $\Phi$ ) and some tested geometries are very close to this point  
10 (see Figure 4 and S values listed in Table 2). It is then naturally expected that rupture transfer  
11 between segments is very easy. It is possible to consider that Segment 1 is not optimally  
12 oriented for failure but is weaker than the surrounding rocks or that Segments 1 and 2 are  
13 governed by different failure conditions. However it is too difficult to quantify such  
14 difference in any general way. By the way, when the Mohr circle is smaller (stress is less  
15 loaded), the rupture transfer possibility is more limited for the fault geometry around the  
16 angle defined by  $\Phi$  and a sequential rupture transfer becomes more dominant. Therefore the  
17 rupture transfer on Segment 2 does not begin before the rupture termination on Segment 1, as  
18 often reported for rupture triggering between segmented faults (e.g. Kase and Kuge, 2001).  
19 This evidence also supports the termination of Segment 1 to the south.

20

21 In the mechanical meaning it may be interesting to study the case where two segments are  
22 separated with a spatial gap. However from the aftershock distribution and the inferred finite  
23 source model, it is natural to think that two faults are overlapping or connecting each other for  
24 the 2007 Niigata-ken Chuetsu-Oki earthquake. In any case, two separated faults have been  
25 already studied by many authors since 1990's (e.g. Harris and Day, 1999; Kase and Kuge,

1 2001). The distance allowing fault jump is often considered a few kilometres, but Kase et al.  
2 (2009) allow a rupture jump over a jog of 10 km distance for the 1668 North Anatolian  
3 earthquake. It is true that the existence of a gap between segments makes it difficult for the  
4 rupture to transfer, but one has to keep in mind that this quantitative possibility strongly  
5 depends on the given stress and frictional condition.

6

7 It is still difficult to compare the detailed processes found in our simulations with the  
8 seismological observations, as different finite source models fit both near-field and  
9 teleseismic data (e.g. Koketsu et al., 2007; Aoi et al., 2008). It is reported that conjugate faults  
10 gives a better fit of the observed seismograms than either one fault model from the view of  
11 statistical analysis using Akaike-Baysian Information Criterion (Horikawa, 2008), but this  
12 does not mean that any characteristic waveform radiated from conjugate faults is identified.  
13 In any cases, the existence of two or three asperities is inferred from north to south along the  
14 fault strike. Each kinematic asperity may briefly correspond to the dynamic rupture of each  
15 segment in our simulations. Comparing our results to those from geodetic inversion  
16 (Nishimura et al., 2008, 2009), the inverted fault slip is concentrated on the area of our fault  
17 model. We cannot currently answer why the 2007 earthquake did not begin on Segment 2 but  
18 on Segment 1. In our simulation, Segment 1 is not always more favourable than Segment 2 in  
19 terms of initial stress and the Mohr-Coulomb criterion (e.g. Figure 4). It is an unsolved  
20 problem whether the rupture starting point (hypocenter) of large earthquakes is selected  
21 deterministically or randomly (Ide and Aochi, 2005). On the particular occasion of the Niigata  
22 region, it is suggested that strengths of highly-dipping faults like the Segment 1 may be  
23 slightly weaker than those of low-dipping faults like the Segment 2, due to a high-fluid  
24 pressures beneath the seismogenic zone and within the fault zones (Sibson, 2007). Spatial

1 distribution of fluid pressure within the crust could be a key factor controlling initiation points  
2 of earthquake ruptures for this region.

3

4 The potential of rupture propagation across the cross-cutting conjugate faults shown by the  
5 present study is significant to seismic hazard analysis. The active fault traces are frequently  
6 used for the estimation of rupture lengths and potential magnitudes. Since the fault traces are  
7 usually not continuous and are composed of several segments separated by fault steps, the  
8 dimension of a fault step is a good indicator for the continuity of the earthquake rupture, and  
9 plays an important role in estimating potential magnitudes. A limiting dimension of fault step  
10 above which earthquake ruptures terminate is reported to range from 3 to 4 km for the case of  
11 strike-slip faults, empirically from the surface trace observations (Wesnousky, 2006).

12 However, about the cross-cutting conjugate thrust-faulting at seismogenic depths, even if the  
13 fault surface traces are separated by distances greater than several ten-kilometres, there  
14 remains a potential for the earthquake rupture to propagate across the cross-cutting conjugate  
15 faults as illustrated in the present study. The rupture transfer across the cross-cutting  
16 conjugate faults as like this example ends by a longer rupture length and a bigger magnitude  
17 than the estimation simply based on each fault surface trace (no interaction is considered  
18 between conjugate faults for any seismic hazard evaluation). It has been observed that cross-  
19 cutting conjugate faults are well developed in fold and thrust zones, where thick sedimentary  
20 layers are piled up, such as the Coalinga region in California (e.g., Eberhart-Philips, 1989)  
21 and the EI Asnam fault zone in Algeria (Chiarabba et al., 1997). Thus, imaging of deep  
22 geometries of active faults beneath the thick sediments is essentially important to evaluate  
23 potential magnitudes within active fold and thrust zones.

24

## 6. Summary

In this study we are able to simulate the dynamic rupture propagation process in a crosscutting conjugate fault system as one proposed scenario of the 2007 Niigata-ken Chuetsu-Oki earthquake. The rupture transfer possibility is very sensitive for the given frictional and stress condition. However for a dip of around 30-45° of Segment 2, its possibility is numerically shown for any frictional levels and independently of the crosscutting distance. Two types of rupture mode are found. If stress is loaded highly under the hydrostatic condition, a simultaneous rupture on both segments is possible along the overlapping part and the triggered rupture on Segment 2 becomes dominant. In other cases where regional stress is not so high but pore pressure plays a significant role to govern the rupture criterion (low frictional coefficient case in this study), rupture does not transfer on Segment 2 until it terminates on Segment 1 regardless of the crosscutting distance. From the view point of the resultant slip amount expected for this dimension of fault planes, a low frictional condition under the high pore pressure existence is strongly preferred. The 2007 Niigata-ken Chuetsu-Oki earthquake offers an important insight into potential scenario of earthquake rupturing, i.e. that they are dynamically governed by the existing fault structure associated with Miocene rifting during an opening stage of the Japan Sea and the subsequent shortening of the crust. The numerical work undertaken in this study should allow us to understand better the mechanics of the dynamic rupture of earthquakes, the importance of structural geology and, consequently, the details of the rupture process.

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8 Tsunamis 2005 and Risques Naturels 2008.

9

## 10 **References**

- 11 [1] Andrews, D. J. (1989), Mechanics of fault junctions, *J. Geophys. Res.*, 94, 9389-9397.
- 12 [2] Aochi, H. and E. Fukuyama (2002), Three-dimensional nonplanar simulation of the  
13 1992 Landers earthquake, *J. Geophys. Res.*, 107, doi:10.1029/2000JB000061.
- 14 [3] Aochi, H., E. Fukuyama, E. and M. Matsu'ura (2000), Spontaneous rupture  
15 propagation on a non-planar fault in 3D elastic medium, *Pure appl. Geophys.*, 157,  
16 2003-2027.
- 17 [4] Aochi H., R. Madariaga and E. Fukuyama (2002), E. Effect of normal stress during  
18 rupture propagation along nonplanar faults, *J. Geophys. Res.*, 107,  
19 doi:10.1029/2001JB000500.
- 20 [5] Aoki, T., M. Furuya, and T. Kato (2008), Coseismic deformation due to the 2007  
21 Chuetsu-oki earthquake (Mw = 6.8), *Earth Planets Space*, 60, 1075-1080.
- 22 [6] Aoi, S., H. Sekiguchi, N. Morikawa and T. Kunugi (2008), Source process of the 2007  
23 Niigata-ken Chuestu-oki earthquake derived from near-fault strong motion data, *Earth*  
24 *Planets Space*, 60, 1131-1135.

- 1 [7] Byerlee, J. D. (1978), Friction of rocks, *Tectonophysics*, 9, 475-486.
- 2 [8] Carena, S. and Suppe, J. (2002), Three-dimensional imaging of active structures using  
3 earthquake aftershocks: the Northridge thrust, California, *J. Struc. Geol.*, 24, 887-904.
- 4 [9] Chiarabba, C., A. Amato, and M. Meghraoui (1997), Tomographic images of the EI  
5 Asnam fault zone and the evolution of a seismogenic thrust-related fold, *J. Geophys.*  
6 *Res.*, 102, 24485-24498.
- 7 [10] Das, S. and K. Aki (1977), A numerical study of two-dimensional spontaneous  
8 rupture propagation, *Geophys. J. R. Astron. Soc.* 50, 643-668.
- 9 [11] Eberhart-Philips, D. (1989), Active Faulting and Deformation of the Coalinga  
10 Anticline as Interpreted From Three-Dimensional Velocity Structure and Seismicity, *J.*  
11 *Geophys. Res.*, 94, 15565-15586.
- 12 [12] Fukuyama, E. (2003), Earthquake dynamic rupture and stress field around the  
13 fault, *J. Geography*, 112, 850-856.
- 14 [13] Fukuyama, E., W. L. Ellsworth, F. Waldhauser, and A. Kubo (2003), Detailed  
15 fault structure of the 2000 western Tottori, Japan, earthquake sequence, *Bull. Seism.*  
16 *Soc. Am.*, 93, 1468-1478.
- 17 [14] Harris, R. A. and S. M. Day (1999), Dynamics of fault interaction – parallel  
18 strike-slip faults, *Geophys. Res. Lett.*, 26, 2089-2092.
- 19 [15] Horikawa, H. (2008), Source model of the 2007 Chuestsu-Oki earthquake,  
20 technical note in Active Fault Research Center News, 75, 1-3.
- 21 [16] Horikawa, H. (2001), Earthquake doublet in Kagoshima, Japan: Rupture of  
22 asperities in a stress shadow, *Bull. Seism. Soc. Am.*, 91, 112-127.
- 23 [17] Ida, Y. (1972), Cohesive force across the tip of a longitudinal-shear crack and  
24 Griffith's specific surface energy, *J. Geophys. Res.*, 77, 3796-3805.

- 1 [18] Ide, S. and H. Aochi (2005), Earthquakes as multiscale dynamic ruptures with  
2 heterogeneous fracture surface energy, *J. Geophys. Res.*, 110, B11303,  
3 doi:10.1029/2004JB003591.
- 4 [19] Ide, S. and M. Takeo (1997), Determination of constitutive relations of fault  
5 slip based on seismic wave analysis, *J. Geophys. Res.*, 102, 27379-27391.
- 6 [20] Kame, N., J. R. Rice and R. Dmowska (2003), Effect of pre-stress state and  
7 rupture velocity on dynamic fault branching, *J. Geophys. Res.*, 108,  
8 doi:1029/2002JB002189.
- 9 [21] Kame, N. and T. Yamashita (1999), Simulation of the spontaneous growth of a  
10 dynamic crack without constraints on the crack tip path, *Geophys. J. Int.*, 139, 345-  
11 358.
- 12 [22] Kase, Y. and K. Kuge (2001), Rupture propagation beyond fault  
13 discontinuities: significance of fault strike and location, *Geophys. J. Int.*, 147, 330-342.
- 14 [23] Kase, Y., H. Kondo and Ö. Emre (2009), Dynamic rupture process of the great  
15 1668 North Anatolian earthquake, *in the proceedings of the French-Japanese*  
16 *workshop on earthquake source*, October 2009, Paris-Orléans, France.
- 17 [24] Kato, A., S. Sakai, N. Hirata, E. Kurashimo, T. Iidaka, T. Iwasaki, and T.  
18 Kanazawa (2006), Imaging the seismic structure and stress field in the source region  
19 of the 2004 mid-Niigata Prefecture earthquake: Structural zones of weakness and  
20 seismogenic stress concentration by ductile flow, *J. Geophys. Res.*, 111, B08308,  
21 doi:10.1029/ 2005JB004016.
- 22 [25] Kato, A., S. Sakai, E. Kurashimo, T. Igarashi, T. Iidaka, N. Hirata, T. Iwasaki,  
23 T. Kanazawa, and Group for the aftershock observations of the 2007 Niigatoken  
24 Chuetsu-oki Earthquake (2008), Imaging heterogeneous velocity structures and

- 1 complex aftershock distributions in the source region of the 2007 Niigataken Chuetsu-  
2 oki Earthquake by a dense seismic observation, *Earth Planets Space*, 60, 1111-1116.
- 3 [26] Kato, A., E. Kurashimo, T. Igarashi, S. Sakai, T. Iidaka, M. Shinohara, T.  
4 Kanazawa, T. Yamada, N. Hirata, and T. Iwasaki (2009), Reactivation of ancient rift  
5 systems triggers devastating intraplate earthquakes, *Geophys. Res. Lett.*, 36, L05301,  
6 doi:10.1029/2008GL036450.
- 7 [27] King, G. and J. Nabelek (1985), Role of fault bends in the initiation and  
8 termination of earthquake rupture, *Science*, 228, 984-987.
- 9 [28] Koketsu, K., H. Miyake and K. Hikima (2007), Source Inversion for the 2007  
10 Chuetsu-oki, Japan, Earthquake: A Case of Difficulty Determining the Source Fault  
11 Plane, *Eos Trans. AGU*, 88(52), Fall Meet. Suppl., Abstract, S54A-05.
- 12 [29] Madariaga, R. and K. B. Olsen (2000), Criticality of rupture dynamics in 3-D,  
13 *Pure Appl. Geophys.*, 157, 1981-2001.
- 14 [30] Matsu'ura, M., H. Kataoka and B. Shibazali (1992), Slip-dependent friction  
15 law and nucleation processes in earthquake rupture, *Tectonophysics*, 211, 135-148.
- 16 [31] Miller, S. A., C. Collettini, L. Chiaraluce, M. Cocco, M. Barchi, and B. Kaus  
17 (2004), Aftershocks driven by a high pressure CO<sub>2</sub> source at depth, *Nature*, 427, 724-  
18 727.
- 19 [32] Nishimura, T., M. Tobita and M. Murakami (2008), Fault model of 2007  
20 M=6.8 Chuetsu-oki earthquake, Central Japan, constructed using geodetic data, *Earth*  
21 *Planets Space*, 60, 1093-1098.
- 22 [33] Nishimura, T., M. Tobita, M. Murakami, T. Kanazawa, and M. Shinohara.  
23 (2009): Fault Model of 2007 M = 6.8 Chuetsu-oki earthquake, central Japan,  
24 constructed using geodetic data, *Advances in Geosciences* (Proceedings of Asia  
25 Oceania Geosciences Society 2008), in press.

- 1 [34] Okamura, Y., M. Watanabe, R. Morijiri, and M. Satoh (1995), Rifting and  
2 basin inversion in the eastern margin of the Japan Sea, *The Island Arc*, 4, 166–81.
- 3 [35] Olsen, K. B., R. Madariaga and R. J. Archuleta (1997), Three-dimensional  
4 Dynamic simulation of the 1992 Landers earthquake, *Science*, 278, 834-837.
- 5 [36] Palmer, A. C. and J. R. Rice (1973), The growth of slip surface in the  
6 progressive failure of over-consolidated clay, *Proc. R. Soc. London, Ser. A.*, 332, 527-  
7 548.
- 8 [37] Scholz C. H. (2002), *The Mechanics of Earthquakes and Faulting*, Cambridge  
9 University Press, New York.
- 10 [38] Shinohara, M., Kanazawa, T., Yamada, T., Nakahigashi, K., Sakai, S., Hino, R.,  
11 Murai, Y., Yamazaki, A., Obana, K., Ito, Y., Iwakiri, K., Miura, R., Machida, Y.,  
12 Mochizuki, K., Uehira, K., Tahara, M., Kuwano, A., Amamiya, S., Kodaira, S.,  
13 Takanami, T., Kaneda, Y., and Iwasaki, T. (2008), Precise aftershock distribution of  
14 the 2007 Chuetsu-oki Earthquake obtained by using an ocean bottom seismometer  
15 network, *Earth Planets Space*, 60, 1121-1126.
- 16 [39] Sibson, R. H (2007), An episode of fault-valve behaviour during compressional  
17 inversion? — The 2004 MJ6.8 Mid-Niigata Prefecture, Japan, earthquake sequence,  
18 *Earth Planet. Sci. Lett.*, 257, 188-199.
- 19 [40] Tada, T. and Yamashita, T. (1997), Non-hypersingular boundary integral  
20 equations for two-dimensional non-planar crack analysis, *Geophys. J. Int.*, 130, 269-  
21 282.
- 22 [41] Takenaka, H., Y. Yamamoto, H. Yamasaki (2009), Rupture process at the  
23 beginning of the 2007 Chuetsu-oki, Niigata, Japan, earthquake, *Earth Planet. Space*,  
24 61, 279-283.

1 [42] Terakawa, T. A. Zoporowski, B. Galvan, and S. A. Miller, Identificatin of the  
2 high pressure fluid source driving the L'Aquila earthquake sequence, EOS, U12A-06,  
3 2009.

4 [43] Townend, J., and M. D. Zoback (2006), Stress, strain, and mountain building in  
5 central Japan, J. Geophys. Res., 111, B03411, doi:10.1029/2005JB003759.

6 Wesnousky, G. S. (2006), Predicting the endpoints of earthquake ruptures, *Nature*,  
7 444, 358-360, doi:10.1038.

8

9

1 Table 1: Model parameters used in this study. Variables are shown in *italic*.

2

<b>Model parameters</b>	<b>Quantity [unit]</b>
P- <i>and</i> S- wave velocities	6000, 3464 [m/s]
Medium density	2500 [kg/m <sup>3</sup> ]
BIEM element size	300 [m]
BIEM time step	0.025 [s]
Segment 1: Dimension (strike, dip, rake)	10.4 x 8 [km <sup>2</sup> ] (215°, 50°, 90°)
Segment 2: Dimension (strike, dip $\delta$ , rake)	20 x 14.4 [km <sup>2</sup> ] (35°, 30,35,40,45,50°, 90°)
Crosscutting distance L	0, 4, 8 [km]

3

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1 Table 2: Initial conditions and frictional parameters assumed in the simulation. The condition  
 2 for Segment 2 is shown for a dip of 30°, 40° and 50°, for references.  $\sigma_3$  is the minimum  
 3 principal stress (vertical direction) extracting the hydrostatic pore pressure (Equation (6)).  
 4

Parameters	Values (in case of stress, unit in MPa)		
Static frictional coefficient $\mu_s$	0.7	0.5	0.3
Effective vertical stress $\sigma_3$ ≡ hydrostatic at 8 km depth	117.6	117.6	117.6
Maximum principal stress $\sigma_1$	433.8	307.9	212.4
Initial shear stress on Segment 1	155.7	93.7	46.7
Initial normal stress on Segment 1	303.2	229.3	173.3
Required strength excess on Segment 1	56.5	20.94	5.28
S value	1	1	1
Dynamic frictional coefficient $\mu_d$	0.33	0.32	0.24
Initial shear stress on Segment 2	-136.9/-155.7/-155.7	-82.4/-93.7/-93.7	-41.1/-46.7/-46.7

---

( $\delta = 30^\circ/40^\circ/50^\circ$ )

Initial normal stress on      196.7/248.3/303.2      165.2/196.2/229.3      141.3/156.8/173.3

Segment 2

( $\delta = 30^\circ/40^\circ/50^\circ$ )

S value on Segment 2      0.01/0.24/1      0.006/0.14/1      0.18/0.037/1

( $\delta = 30^\circ/40^\circ/50^\circ$ )

---

1

1 Figure captions

2 Figure 1: The 2007 Niigata-ken Chuetsu-Oki, Japan, earthquake. Relocated aftershock  
3 distributions along cross sections and the inferred faulting model as shown by gray-thick lines  
4 with arrows denoting slip directions (after Kato et al., 2008). Aftershocks are colored to the  
5 depths. The Moment tensor for the mainshock (the NIED) clearly shows reverse faulting but  
6 the aftershock distributions demonstrate north-west dipping fault in the north and south-east  
7 dipping fault in the south. The direction of the maximum principal stress axis  $\sigma_1$  is drawn as  
8 thick-white arrows.

9

10 Figure 2: Fault geometry with respect to the aftershock distribution in three-dimensional  
11 space from different angles of view (cross section view in the left panel, and 3D view from  
12 southern-east-upper direction in the right panel).The variables in the simulations are the dip  
13 angle ( $\delta$ ) of Segment 2 (SE dipping in the south) and the overlapping length (L) between two  
14 segments. The mainshock hypocenter is located on the deep part of the Segment 1.

15

16 Figure3: Slip-weakening law defined by Equation (1).

17

18 Figure 4: Coulomb rupture criterion and Mohr circle in the case of  $\mu_s$  (a) 0.7 and (b) 0.3. In  
19 both cases, the hydrostatic condition at 8 km depth is supposed, namely  $\sigma_3 = 117.6$  MPa. The  
20 dynamic stress level  $\mu_d$  is drawn so as to S value to be 1 on Segment 1 (see the amount  
21 signed by the arrows comparing to Figure 3). The initial condition on Segment 1 is  
22 represented by a solid circle and the initial conditions on Segment 2 are shown by triangles  
23 for a dip of 30, 40 and 50°.

24

25

1 Figure 5: Snapshots of rupture propagation. For each simulation, slip velocity, fault slip, shear  
2 stress change and normal stress change are illustrated at different simulation time. The bold  
3 lines show the cross section with the other segment. The arrows show the fault strike in  
4 horizontal axis and dip in vertical axes. (a) Simultaneous rupture transfer on Segment 2. (b)  
5 Sequential rupture transfer from Segment 1 to 2. (c) Case of no rupture transfer.

6

7 Figure 6: Diagram of rupture mode obtained through numerical simulations. (a)  $\mu_s = 0.7$ , (b)  
8  $\mu_s = 0.5$  and (c)  $\mu_s = 0.3$ , respectively. The double circles represent simultaneous rupture on  
9 both segments as shown in Figure 5(a). The single circles means sequential rupture transfer  
10 from Segment 1 to Segment 2 such as Figure 5(b). The crosses are the cases that rupture does  
11 not transfer on Segment 2.

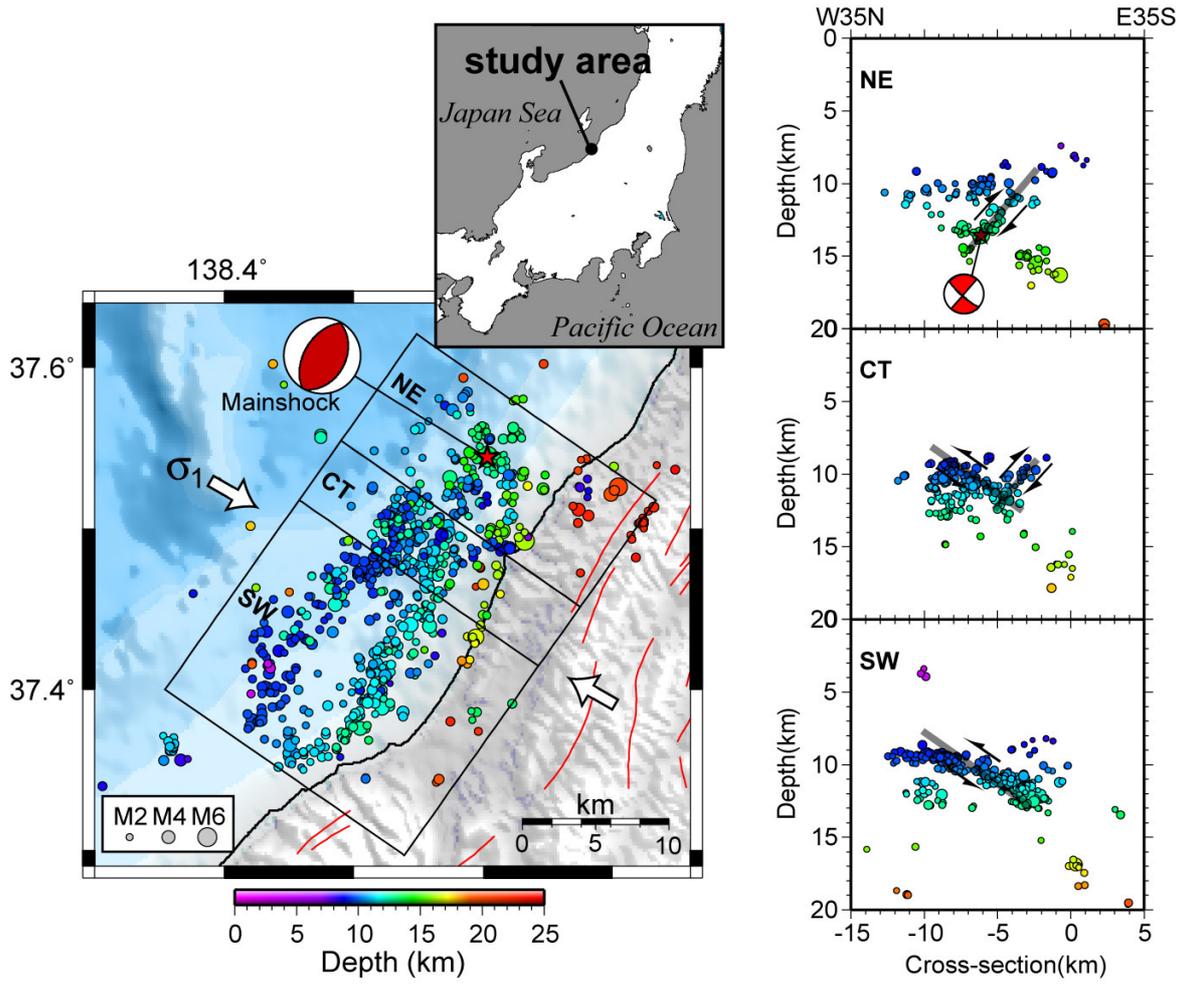
12

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14 Figure 7: Rupture time on Segment 2 when  $\mu_s = 0.3$  and  $\delta = 45^\circ$  for variable over-lapping  
15 lengths  $L$ . Rupture time at each point on Segment 2 is shown between 3 and 8 seconds after  
16 the origin time. The red bold line or point show the position where Segment 1 is cross-cutting.

17

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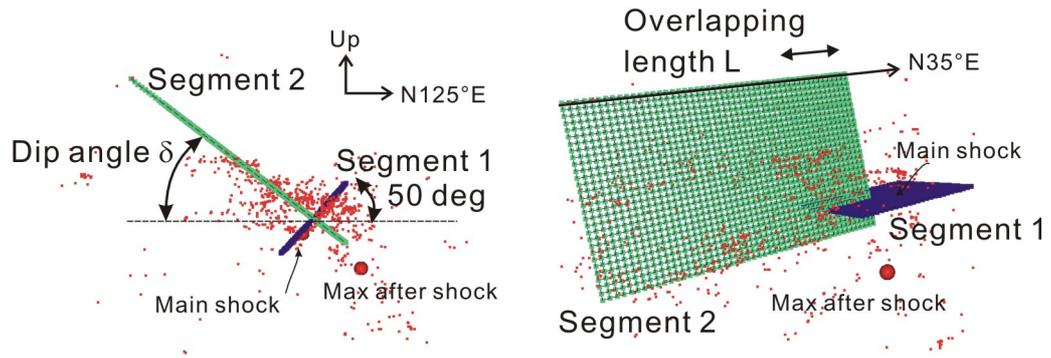


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Figure 1.



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Figure 2.

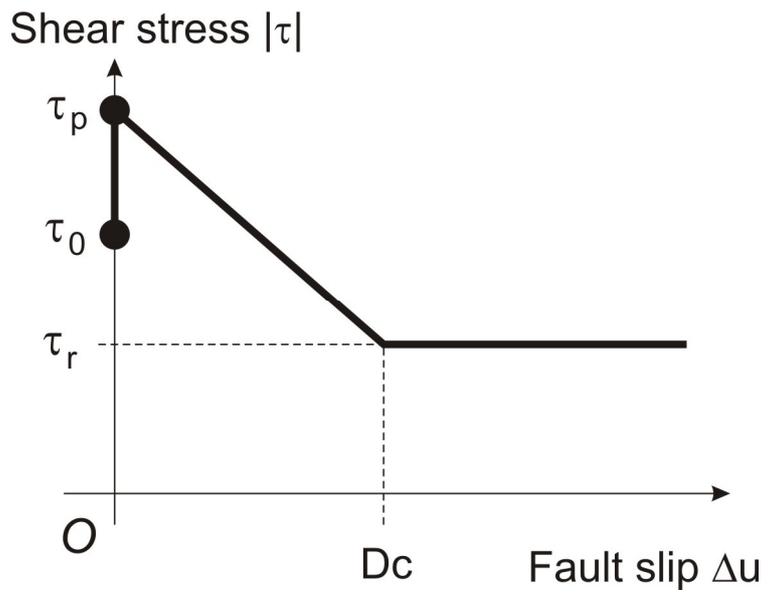
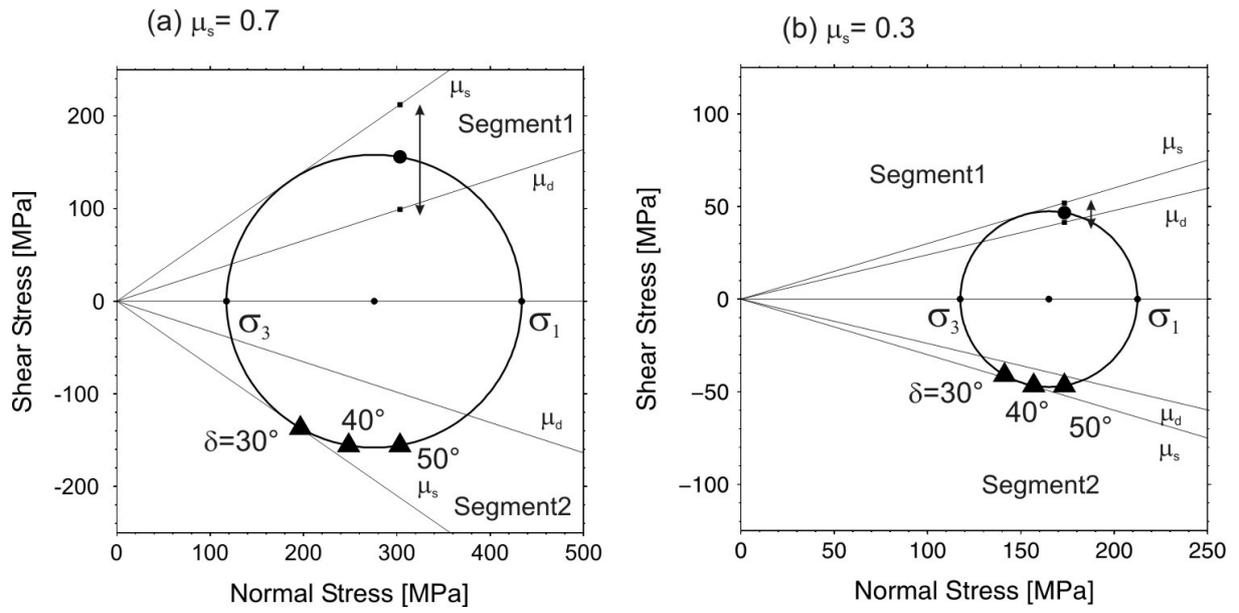


Figure 3.

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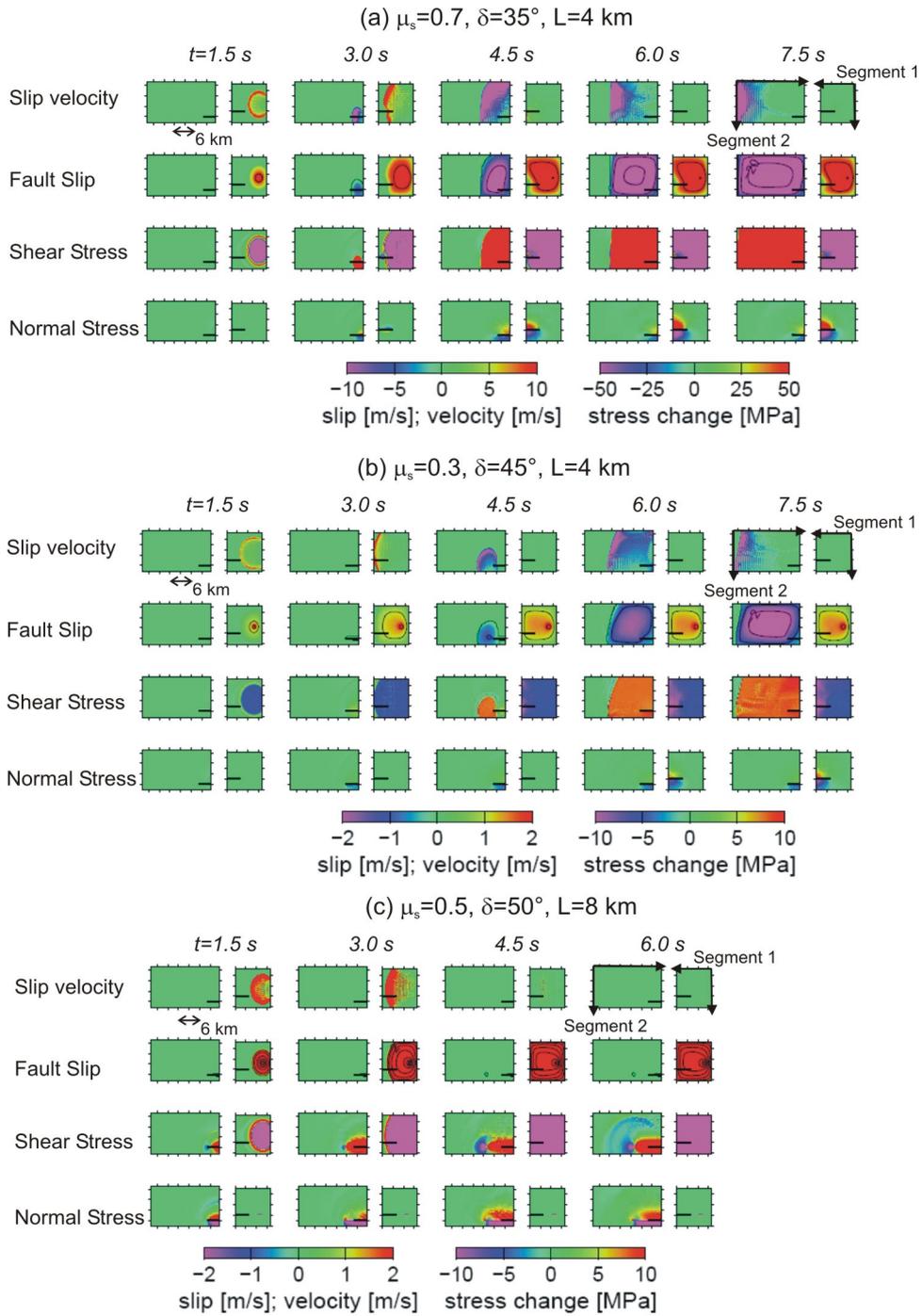


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Figure 4 :

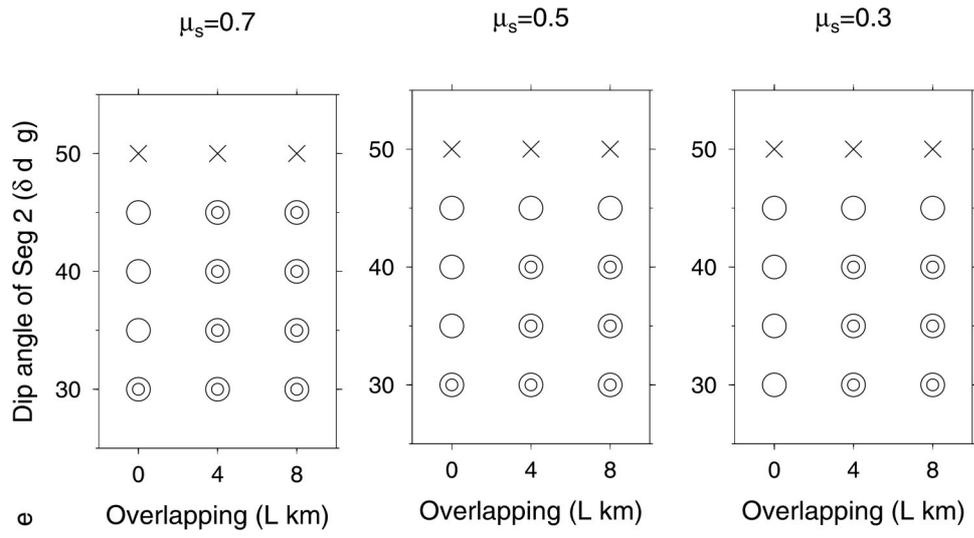
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Figure 5.

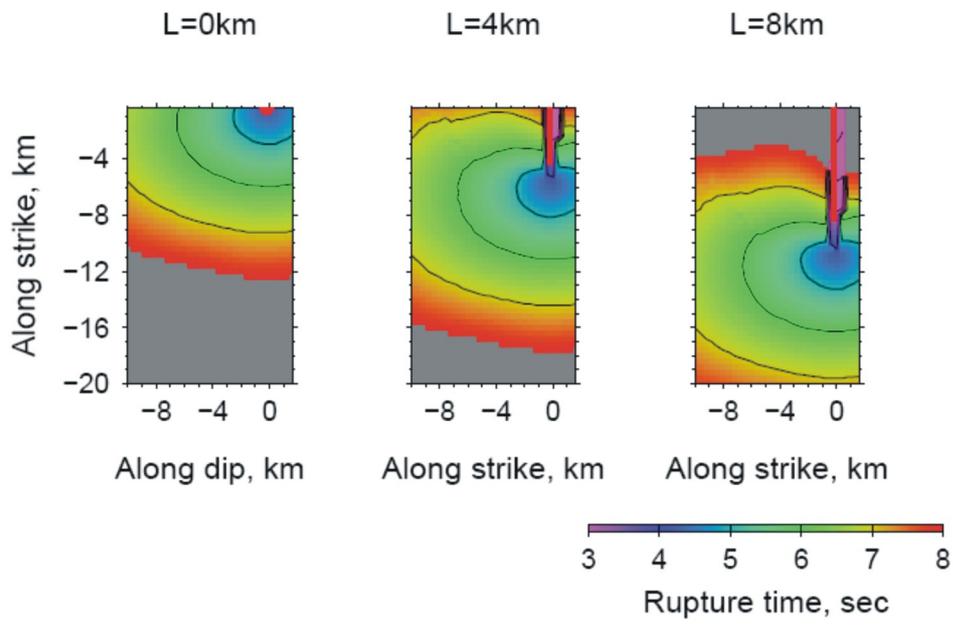


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e Figure 6.

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3 Figure 7.